

Zentralanstalt für Meteorologie und Geodynamik

Emerging technologies in the numerical forecasting: regional ensemble prediction and nowcasting

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With contribution of many colleagues from ZAMG, LACE, ALADIN...

Outline

- Regional ensemble prediction
 - ALADIN-LAEF (Limited Area Ensemble Forecasting)
- Nowcasting
 - INCA (Integrated Nowcasting through Comprehensive Analysis)
- Summary

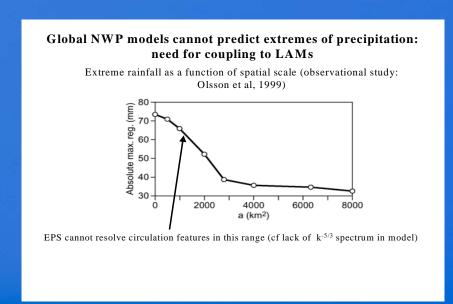


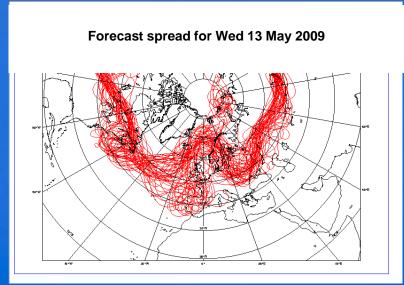
Regional ensemble prediction

Atmosphere: chaotic and highly non-linear nature

Small errors in analysis, model physics can grow rapidly and become large, even in a matter of hours.

Global EPS: lower resolution. medium range



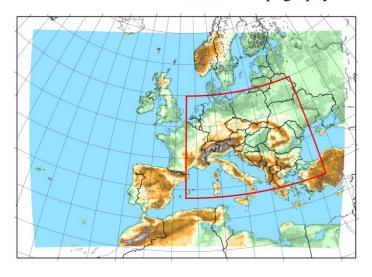




ALADIN-LAEF: Limited Area Ensemble Forecasting

Ensemble size	16+1
Horizontal resolution	18 km
Vertical resolution	37 levels
Runs/Day	2 (00, 12 UTC)
Forecast range	60h
Output-Frequency	1h
Model time step	720s
Coupling-Model	ECMWF-EPS
Coupling-Update	6h

ALADIN-LAEF Domain & Topography



Atmosphere perturbation: Blending ALADIN Bred + ECMWF EPS

Surface perturbation:

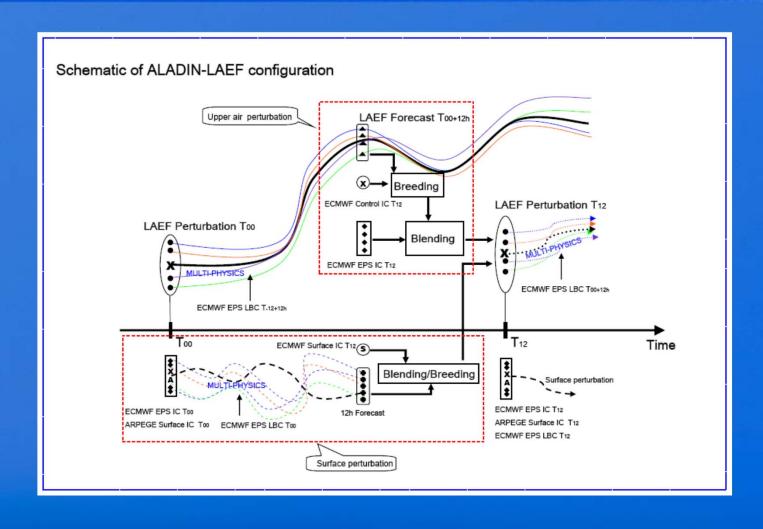
Non-Cycling surface Breeding

Model perturbation: multi-physics

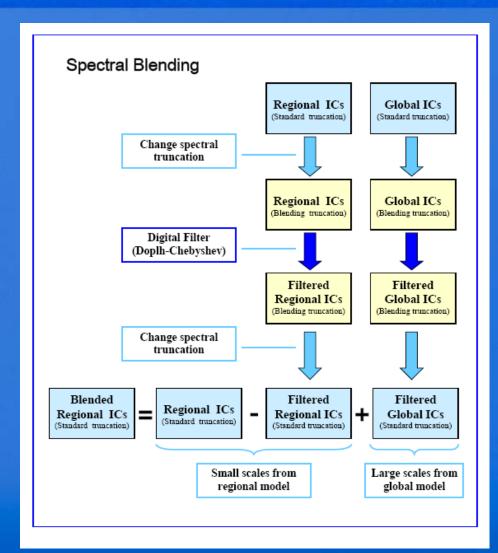


Operation at ECMWF, SMS suite, Time Critical Application II 2m Minimum Temperature [°C], ENS-Mean, 20070517, 12 UTC + 54 **ARPEGE ECMWF** analysis **EPS fcst ECMWF MARS ECMWF/HPC ECMWF/HPC** Data premodel **ECMWF/HPC** processing integration data postprocessing **ECMWF MARS** LACE: CZ, **ZAMG** SLO, SK visualisation & products Zentralanstalt für Meteorologie und Geodynamik

ALADIN-LAEF: design



Blending: Theory



The Blending method

Following the idea by Machenhauer and Haugen (1987) and considering perturbed variable G from global ensemble, and R from regional ensemble, both G and R are valid at the resolution of regional model, their full harmonic Fourier expansions are

$$G(x) = \sum_{m=1}^{M} G_m e^{ik\pi(m\omega/L_n)}$$
; $R(x) = \sum_{m=1}^{M} R_m e^{ik\pi(m\omega/L_n)}$ (1)

where M is the maximum wave numbers, and L_x is the horizontal wavelengths in x direction. Let J be the number of grid-points of the whole regional domain along the x direction, the inverse truncated Fourier transform. Let the spectral coefficients G_m and R_m are obtained by:

$$G_m = \frac{1}{J} \sum_{j=0}^{J-1} G(j) e^{-2i\pi(jmiJ)}$$
; $R_m = \frac{1}{J} \sum_{j=0}^{J-1} R(j) e^{-2i\pi(jmiJ)}$

As we mentioned above, the blending is applied on full grid-point resolution of the regional model but with a lower spectral resolution (we call it the blending truncation), which represents the scale resolved by the global data assimilation. We change the spectral fruncation from the original resolution to the blending spectral resolution, For both G and R, their full harmonic Fourier expansions are given by:

$$G(x) = \sum_{m=-LM}^{LM} G_m^{LOW} e^{iLe(m\pi L_s)}$$
; $R(x) = \sum_{m=-LM}^{LM} R_m^{LOW} e^{iLe(m\pi L_s)}$ (3)

where LM is the maximum wave numbers of the blending spectral truncation, the inverse truncated Fourier transform, i.e. the spectral coefficients G_m^{LOW} and R_m^{LOW} are obtained by:

$$G_n^{LOW} = \frac{1}{J} \sum_{i=0}^{J-1} G(j) e^{-2i\pi(jniJ)}$$
; $R_n^{LOW} = \frac{1}{J} \sum_{i=0}^{J-1} R(j) e^{-2i\pi(jniJ)}$

The blending truncation is determined by the resolution of global analysis, i.e. the resolution of the data assimilation of the global mode regional model. The estimation of the blending spec For the low-frequency range, $H(\theta)$ falls from 1 to $r=1/T_{2b}(x_0)$ as $|\theta|$ goes from 0 to θ_0 nomials, which is defined as: obtained as follows:

$$T_R^{cut} = \sqrt[3]{T_G^{\alpha^2} \times T_R^f}$$

Where $T_{\mathcal{Q}}^{a}$ is the spectral truncation of the global data equivalent truncation of global model corresponding to The maximum wave numbers of the blending spectral truncation is then calculated

$$M = M \times (T^f / T^{out})$$

The blending ratio $T_k^f/T_k^{\rm out}$ should be larger than ratio between the average resolution of global model over the regional domain and the resolution of the regional

The scale selection in the blending is by employing the digital filter technique, which was originally used for the initialization of meteorological fields in NWP. Its detailed description can be found in Lynch and Huang (1992).

For any model state G_m^{LOW} and R_m^{LOW} , denoted in the following as f_n , known at $\{f, 1, \dots, f_n, f_n, f_n, f_n, f_n\}$, it may be regarded as the Fourier coefficients of a function

$$F(\theta) = \sum_{i}^{mN} f_n \times e^{-im\theta}$$
(7)

where θ is the digital frequency. The filtering of f could be conducted by multiplying

$$H(\theta) = \begin{cases} 1 & \text{if } |\theta| \le \theta_c \\ 0 & \text{if } |\theta| \ge \theta_c \end{cases}$$
(8)

 θ_c is the *cutoff frequency*. In ALADIN digital filter initialization, the non-recursive Dolph-Chebyshev filter is applied (Lynch et al. 1997), which is given by

$$\frac{I_{2N}(X_0 \cos(\theta / 2))}{I_{2N}(X_0)}$$

h is specified by choosing a

(11)

and for the high frequency range $\theta_0 \le \|\theta\| \le \pi$, $H(\theta)$ oscillates within $\pm r$. Let f_n^* denote the low-frequency part of f_a , clearly:

$$F(\theta) = \sum_{n = N} f_n^* e^{-n\theta}$$
 (12) als can be obtained from the

$$f_s^a = \sum_{k=0}^{k=0} (H \cdot F)_k f_{n-k} = \sum_{k=0}^{k=0} h_k f_{n-k}$$
 (13)

This is a non-recursive digital filter, since f_n^* depends on both past and future values of $f_{\rm s}$, but not other outputs values (Lynch et al 1997). $h_{\rm fl}$ is given by:

$$h_n = \frac{1}{2N+1} \left[1 + 2r \sum_{n=1}^{N} T_{2N}(x_0 \cos \frac{\theta_n}{2}) \cos m \theta_n \right]$$
 (14)

The solution of the model, integrated from $-t_N$ to t_N , is weighted averaged:

$$f^{*}(0) = \sum_{i=1}^{n} h_{ii} f_{ii}$$
 (1)

so that at the end of the DFI a balanced initial state is achieved. In ALADIN DFI, the filter order N is determined by the time step ΔT based on blending truncation and the

Suppose the results of DFI on $G_m^{\ LOW}$ and $R_m^{\ LOW}$ are $G_m^{\ TLOW}$ and $R_m^{\ TLOW}$. We obtain the large scale part of the perturbed ICs of global and regional model by

$$G^{LDP}(x) = \sum_{i}^{M} G_{ii}^{*} G^{ECOF} e^{2i\pi(m_{i}/L_{i})}$$
; $R^{LDP}(x) = \sum_{i}^{M} R_{m}^{*} G^{OF} e^{2i\pi(m_{i}/L_{i})}$ (16)

The symbolic equation of Blending can be summarized after Brožková et al. (2006) and Derkova and Bellus (2007):

$$IC_{blood} = R - R^{l\phi} + G^{l\phi}$$
 (17)

Where IC_{blood} denotes initial condition after blending. After the blending, the filed itself keeps the large scale from global assimilation, and small scale from regional model, but it might be still noisey due to the pure interpolation, therefore an additional DF initialization with a narrow stop-band, could be applied to suppress the noise due to the interplolation.



Blending: feature

Blending global ECMWF EPS with LAEF Breeding

To combine the large-scale uncertainty from ECMWF EPS with the small-scale uncertainty generated by Breeding in LAEF.

LAEF Initial perturbations:

- ➤ The scale of the LAEF perturbation is in accordance with the scales of variability resolved by the model.
- ➤ The LAEF perturbations are consistent with the perturbation coming through the lateral boundary.
- ➤ The LAEF perturbations are effective immediately from the initial time.

Met-office NCEP NCAR



ALADIN-LAEF performance

What is the more added value of LAEF to its counterpart ECMWF EPS?

ECMWF EPS
$$\longleftrightarrow$$
 LAEF

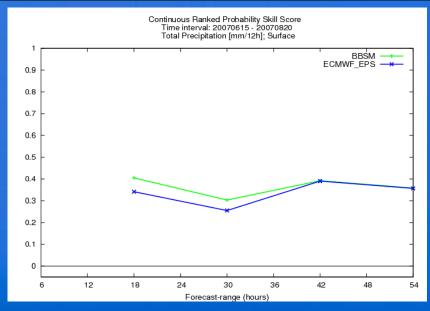
Is LAEF better than its existing high resolution deterministic ALADIN forecast?

ALADIN-Austria \leftarrow LAEF



ALADIN-LAEF vs. ECMWF EPS

	ALADIN-LAEF	ECMWF-EPS
Resolution	18km; 37 Levels	T _L 399; 62 Levels
Ens. Size	16	50
Model	ALADIN	ECMWF-IFS



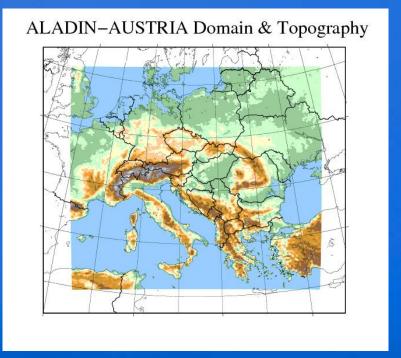


ALADIN-LAEF vs. ALADIN-Austria

Is ALADIN-LAEF adding value to its existing high resolution deterministic ALADIN forecast?

Horizontal resolution	9.6 km
Vertical resolution	60 Levels
Runs/day	00,06,12,18 UTC
Forecast range	72h / 60h
Output- Frequence	1h
Time step	415s
Coupling-Modell	ARPEGE
Coupling-Update	3h

ALADIN-Austria: deterministic





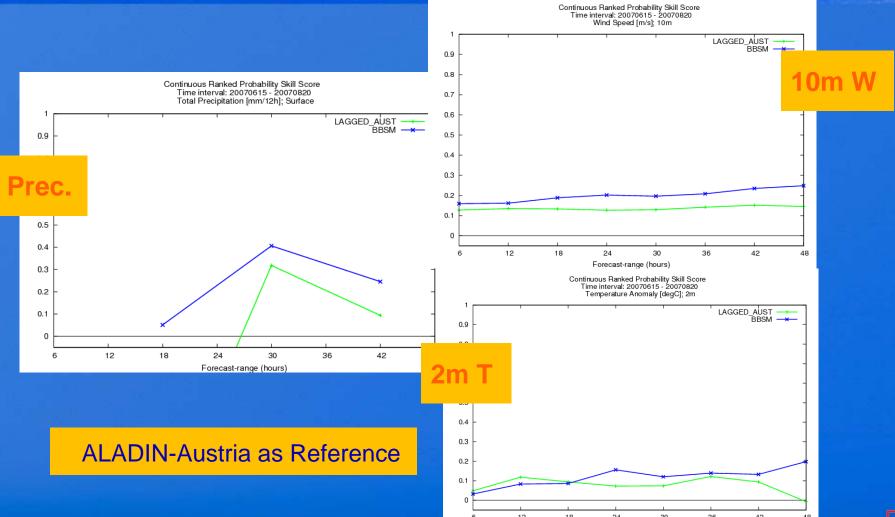
ALADIN-LAEF vs. ALADIN-Austria

	ALADIN-LAEF	ALADIN-AUSTRIA
Resolution	18km;37Levels	9.6km;60 Levels
Ensemble size	16 members	5 members (time lagged)
Forecast	Ensemble mean	deterministic

ALADIN-Austria: time lagged EPS

00 UTC:	00	06	12	18	24	30	36	42	48	54	60	66	72
06 UTC:		00	06	12	18	24	30	36	42	48	54	60	66
12 UTC:			00	06	12	18	24	30	36	42	48	54	60
18 UTC:				00	06	12	18	24	30	36	42	48	54
00 UTC:					00	06	12	18	24	30	36	42	48



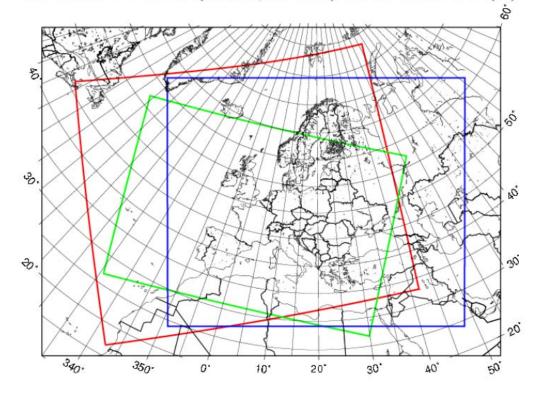


18 24 30 36 42 48
Forecast-range (hours)

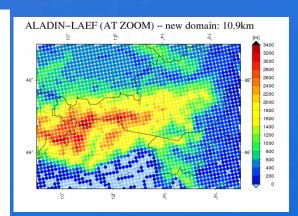
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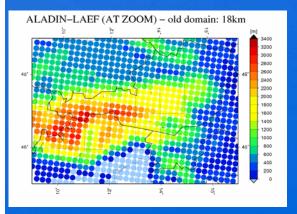
ALADIN-LAEF upgrade, cooperation with Turkey

ALADIN-LAEF (old:G, new:B) vs GLAMEPS (R)



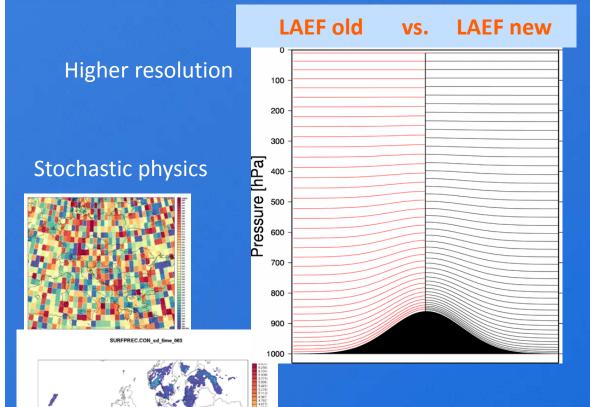
::Fig.01 Domain boundaries of the operational ALADIN-LAEF (green), new redefined ALADIN-LAEF (blue) and GLAMEPS (red).



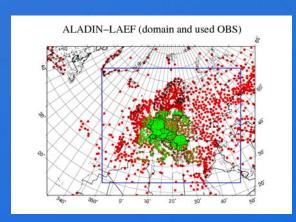




ALADIN-LAEF upgrade, cooperation with Turkey



Optimsied multi-physics



Ensemble land surface assimilation

NAMEUST	MICROPHY TUNING	DEEP TUNING	SHALLOW FUNING	RADIATION TUNING	TURBULEN TUNING	BUST DIAG. TUWNS	SCREENING LEVEL TUNIN DAG.	
MP01	ALARO-KR	3MT	IFG87	ireos	NG06	RAFTUR	NSCREO	
MPGZ	ALARD X8 MEH	3MT 02L, 03H	#G87	#505 R3H, R3H, R4T	F606	RAFTLR/RA FTHE/BAFB	NSCREO/NS CREL/NSCR	
MP03	MANDON MILMAL MR	SMT DZH, DSL	FG87	#505 #11, R31, 847	F606	RAFTUR/RA FTKE/RAFB	NSCREQ/NS CREIL/NSCR	
MP04	ALARO NE MIE	SMT DJH, DSL	FGE	#G05 #11, #31, 847	FG06	RAFTUR/RA FTKE/RAFB	MSCREO/MS CREI/NSCR	
MPGS	ALARO-JER MITH, MOH, MIT	SMT D2H, D3L	#GET	#G05 #31, #31, #47	FGM	FTIE/RAFB	MSCREC/WS CREL/NISCR	
MPGE	ALARO SM	BMT	#G82	FGOS AN	F666 72H,TR	FLAFTURVIKA FTKE/RAFB	NSCREG/NS CREL/NSCR	
MPOT	ALMIO SM	SMT DST	#682	JFG05 AFT	FG06 72L 73H	RAFTUR/RA FTHE/RAFB	NICREO/WS CREIL/NISCR	
MPGS	FORES	pg .	KFB	ECMWF	CBR	RAFTKE	NSCREO	
MP09	LOPEZ	BG DIT	eru.	LOWNE	CBH	RAFILE/EA FTEE/EAFB	ASCRED/WS CHEL/NISCR	
MP10	LOPEZ	es .	418	ECMMF	CBR	RAFTUR/RA FTRE/RAFB	MSCREQ/MS CREL/NISCR	
MP11	LOPEZ	as our	698 52	EOWIF	CNR	FLAFTUR/RA FTKE/RAFB	MSCREQ/WS CREI/NSCR	
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