

# What bout water vapor and condensation at Dome C

*C. Genthon<sup>1</sup>, J.-B. Madeleine<sup>1</sup>, R. Forbes<sup>2</sup>, H. Gallée<sup>3</sup>, E. Vignon<sup>4</sup>, L. Piard<sup>3</sup>, F. Hourdin<sup>1</sup>, F. Lemonnier<sup>1</sup>, M. Casado<sup>5</sup>, G. Camporeale<sup>6</sup>*

*(1) Laboratoire de Météorologie Dynamique, Paris*

*(2) ECMWF, Reading*

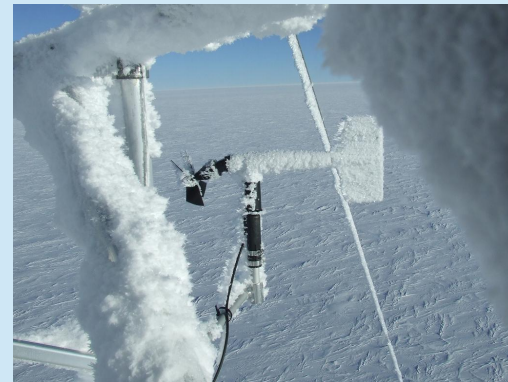
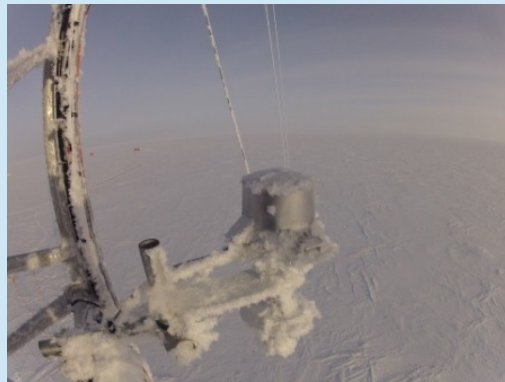
*(3) Institut des Géosciences de l'Environnement, Grenoble*

*(4) Ecole Polytechnique Fédérale, Lausanne*

*(5) Laboratoire des Sciences du Climat et de l'Environnement, Saclay*

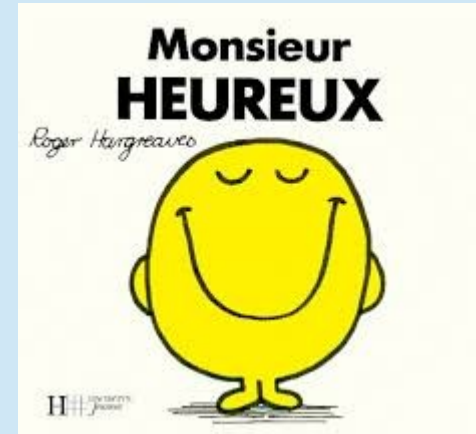
*(6) ENEA, Casaccia*

Little atmospheric moisture ( $T^{\circ} \sim -30 / -70^{\circ}\text{C}$ ), yet a lot of frost => necessarily near or above saturation ?



There used to be a time when climate modeling was simple.

For instance, basic equilibrium thermodynamics would apply:



With good old FORTRAN 77

IF

    partial pressure of water vapor .GT. saturation partial pressure

THEN

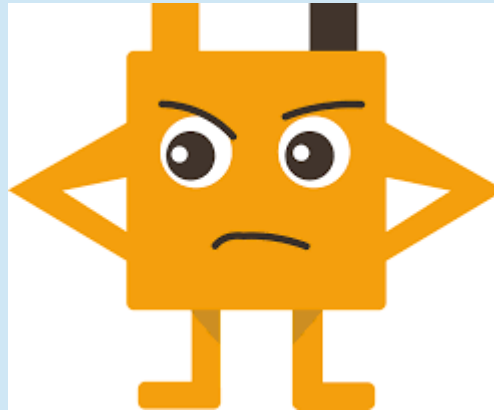
    Dump excess water, make cloud and rain ( $T > 0^{\circ}\text{C}$ ) or snow ( $T < 0^{\circ}\text{C}$ ),

ELSE

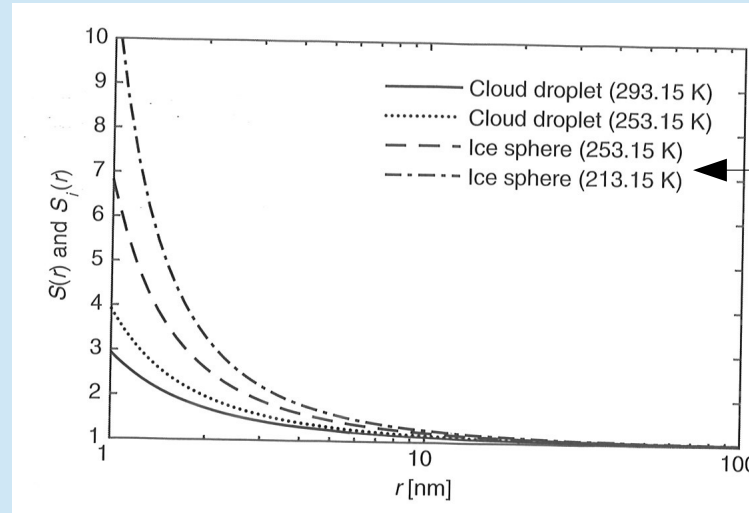
    Clear sky, no precip

END IF

Good old time. Until we found (or accepted) that supersaturation exists in the atmosphere.



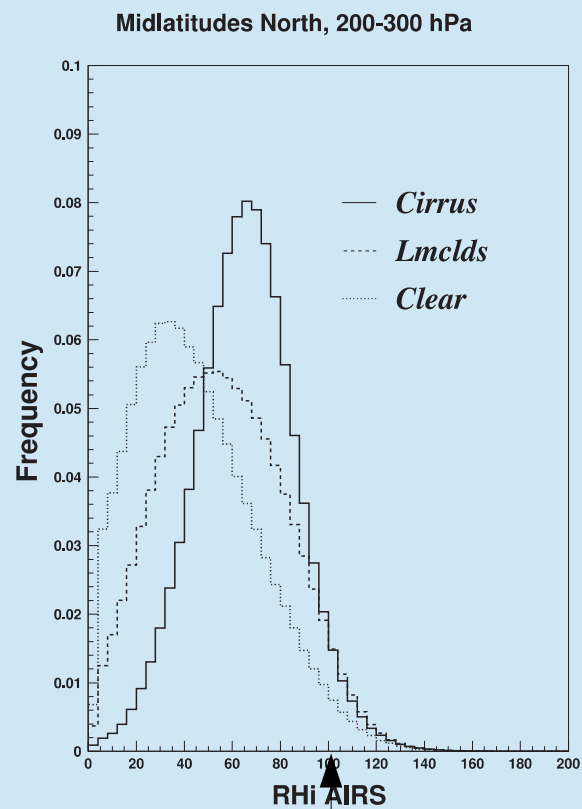
Enhanced thermodynamics: Supersaturation is required for homogeneous nucleation to spherical droplet / ice crystal to occur (Kelvin effect). The required supersaturation is larger as temperature is colder.



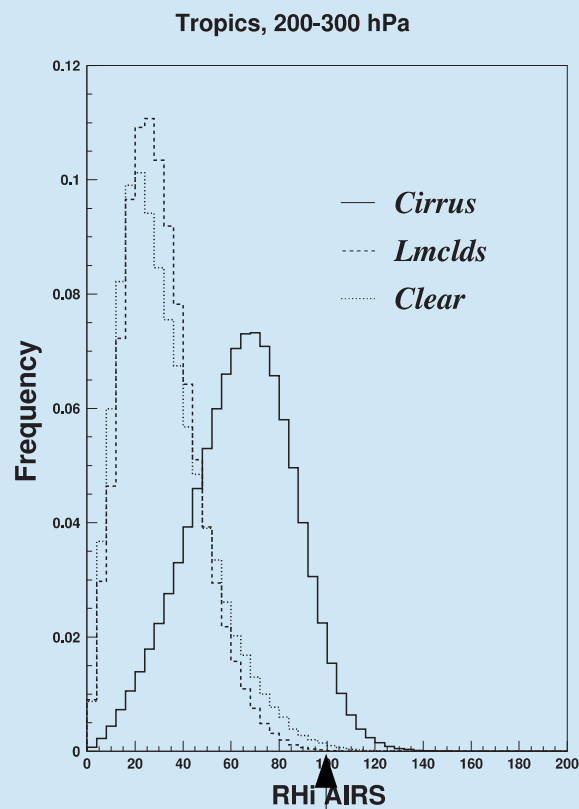
Upper troposphere  
Antarctic surface

*From Lohmann, Lüönd and Mahrt, 2016. An Introduction to clouds, Cambridge University Press, Fig 8.1*

In addition, if supercooled water, then humidity  $> \sim$  supersaturation with respect to liquid  $\Rightarrow$  largely supersaturated with respect to ice



100 %  
 With respect to ice  
 (wri)



100 %  
 With respect to ice  
 (wri)

From Atmospheric InfraRed Sounder (AIRS)

Good old time. Until we found (or accepted) that supersaturation exists in our atmosphere.

- Including supersaturation with respect to ice (wri)
- Not only exists but also is frequent
- Not only in clear sky but also in cloudy conditions

And it must be accounted / parameterized in meteorological and climate models to account for such highly sensitive clouds as the high altitude cirrus

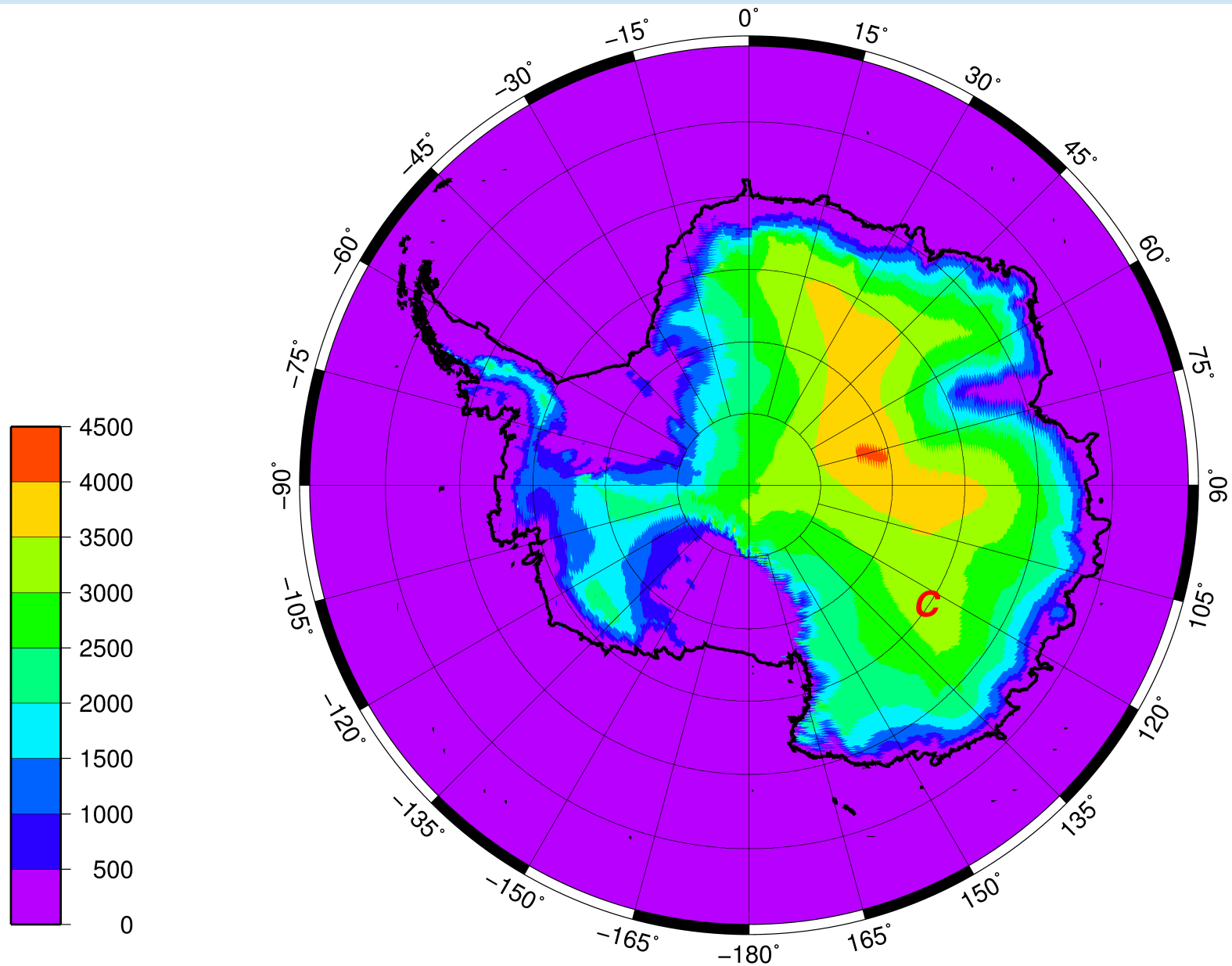
Cold temperature + low CCN / IN => high altitude, not at the surface

=> An issue for high altitude clouds (cirrus) only

=> Obtain in situ observation to verify parameterization is difficult

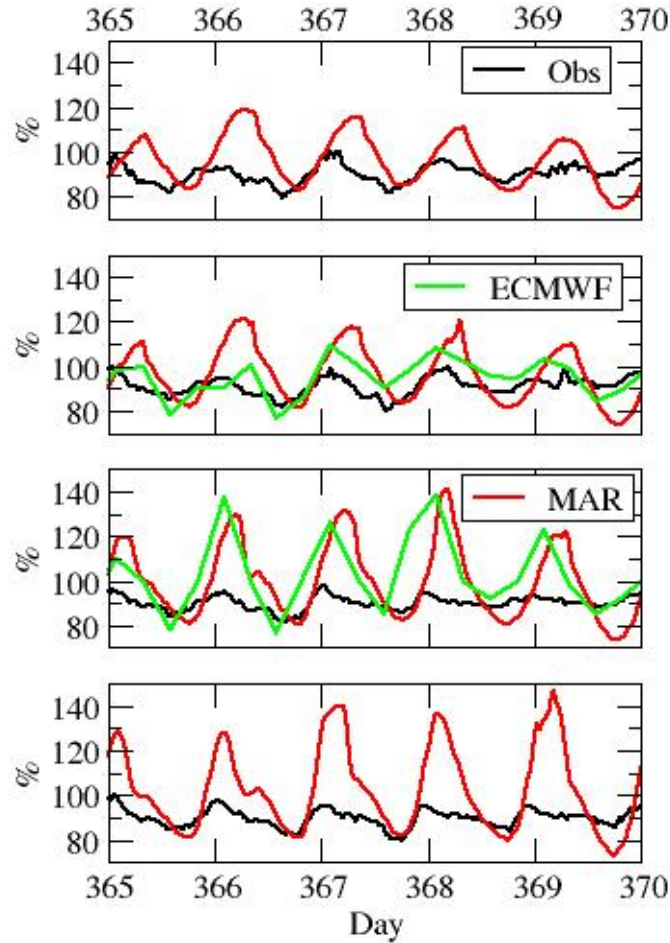
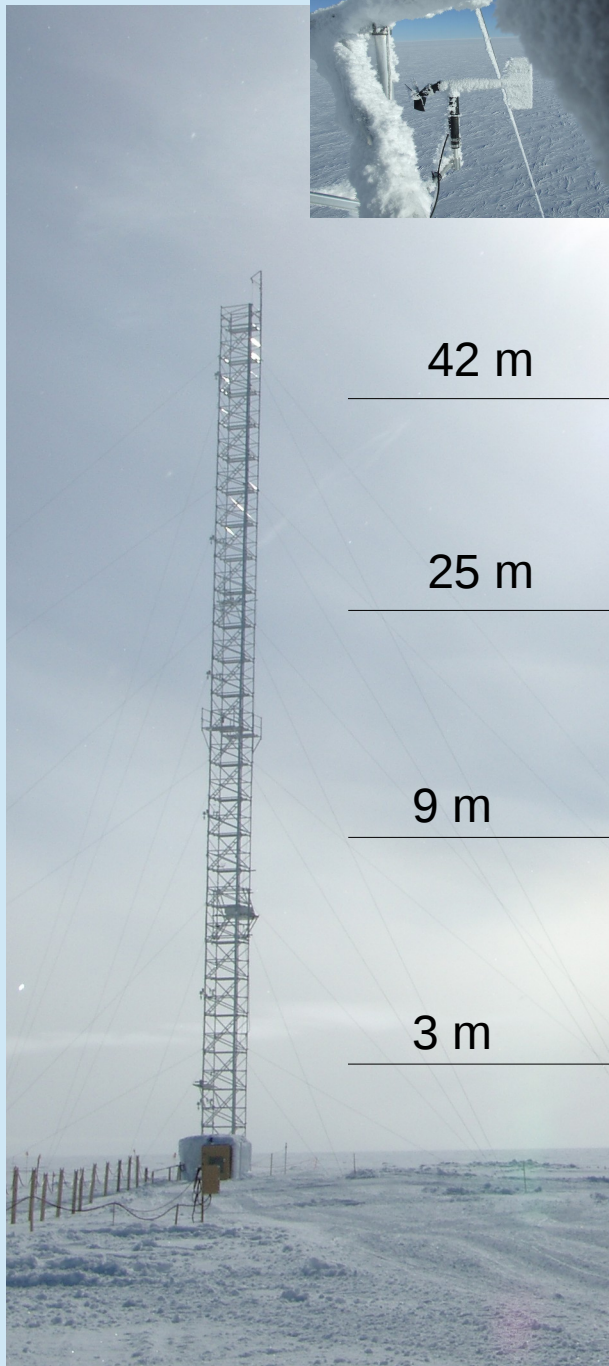
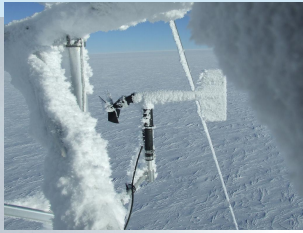
Where do we have cold temperature, low CCN/IN, right at the surface?

Dome C, East Antarctic plateau, 75°060S, 123°200E, 3233 m a.s.l.



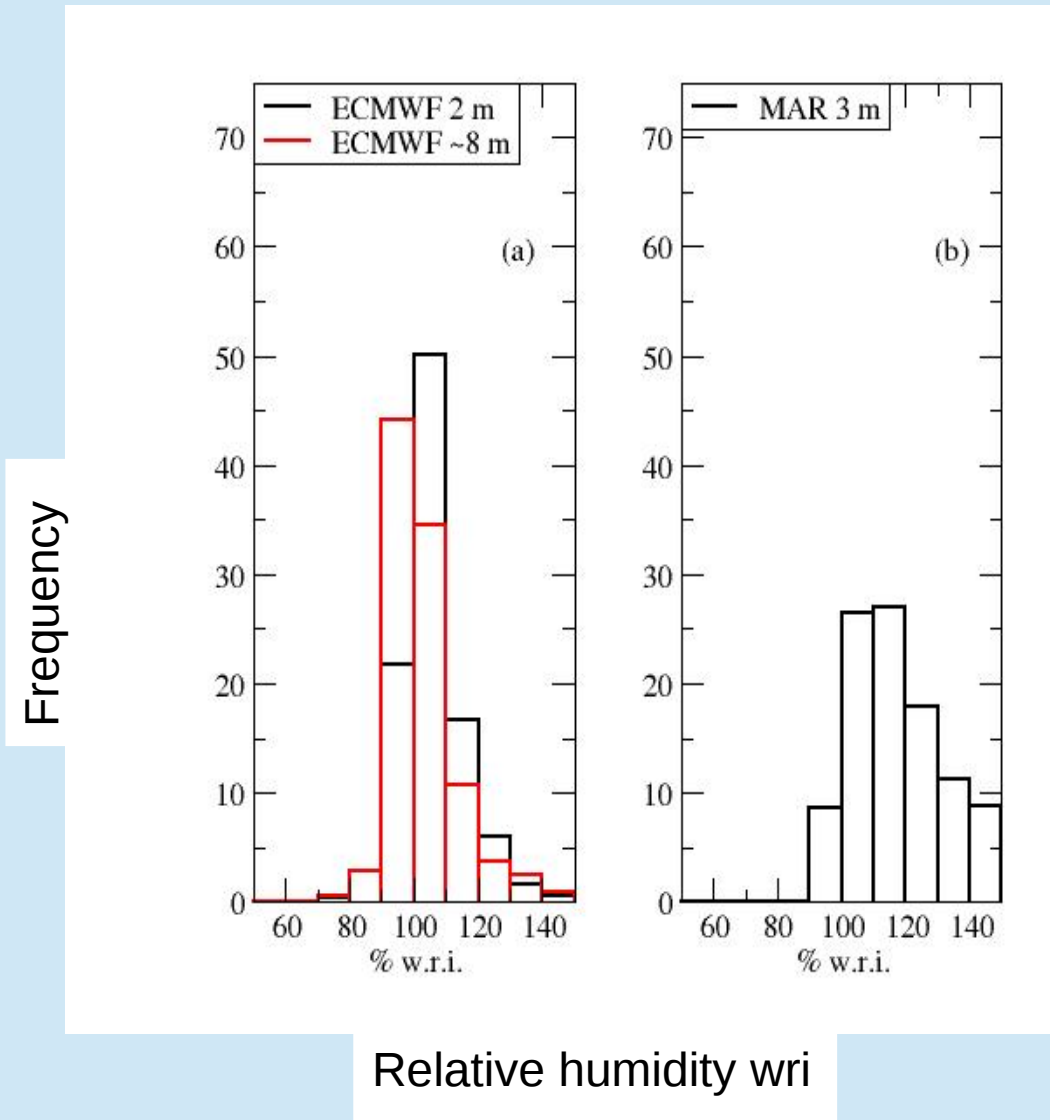


=> Standard (incl. WMO approved) instruments do not report supersaturation



OBS ———  
ECMWF ———  
MAR ———

The models (with cold microphysics) suggest that supersaturation is very frequent



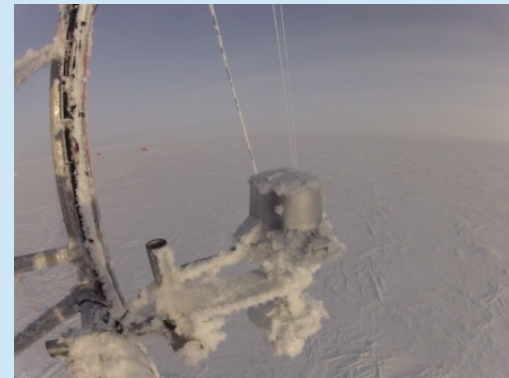
Where moisture is measured

Where frost deposits / excess moisture is dumped



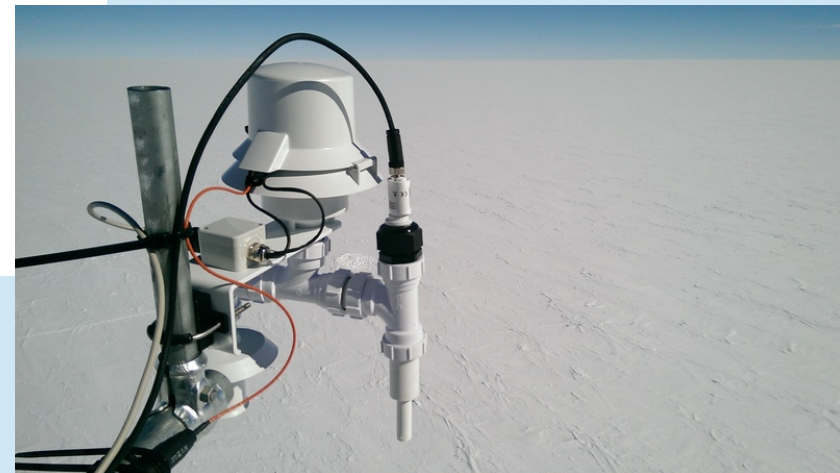
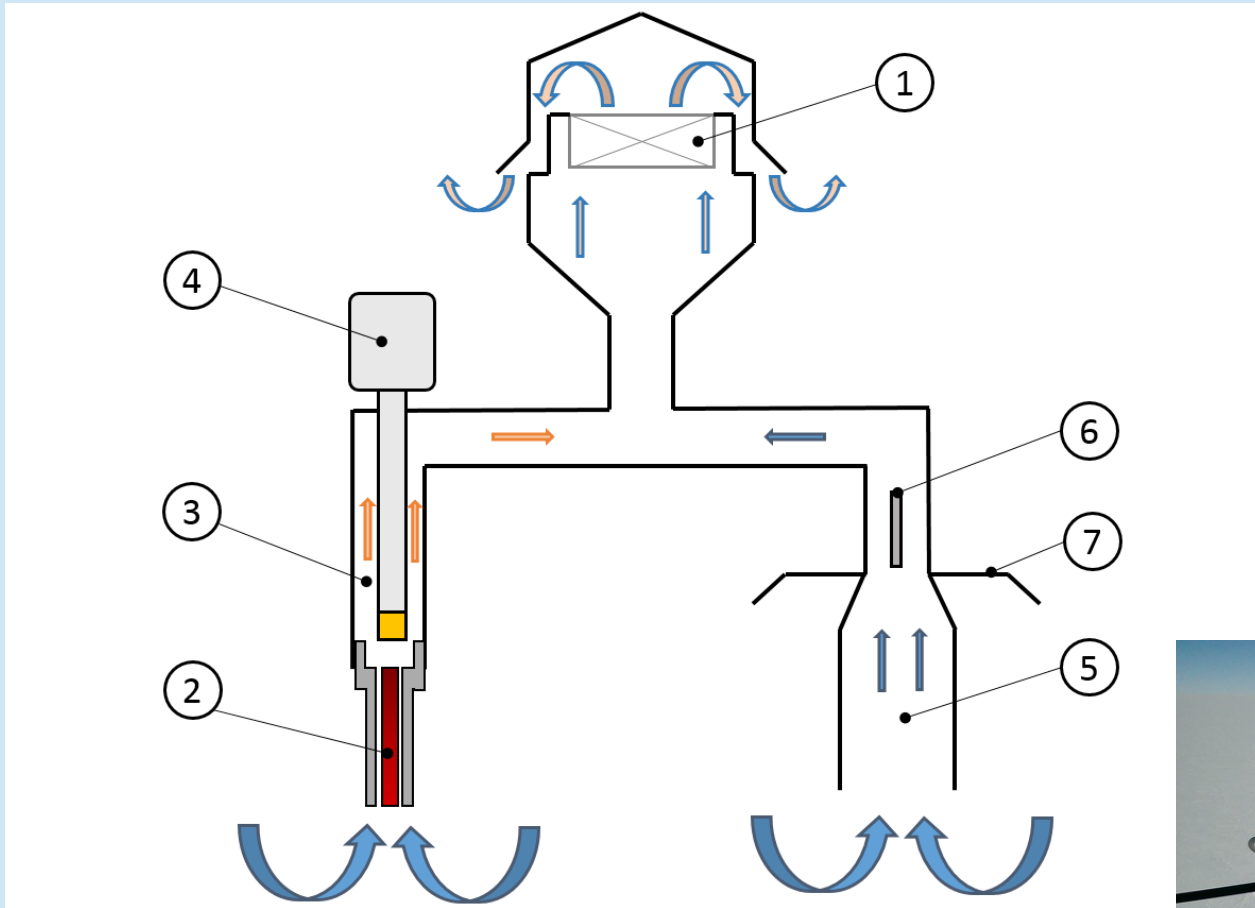
A typical commercial hygrometer (Vaisala HMP155)

In its radiation shield



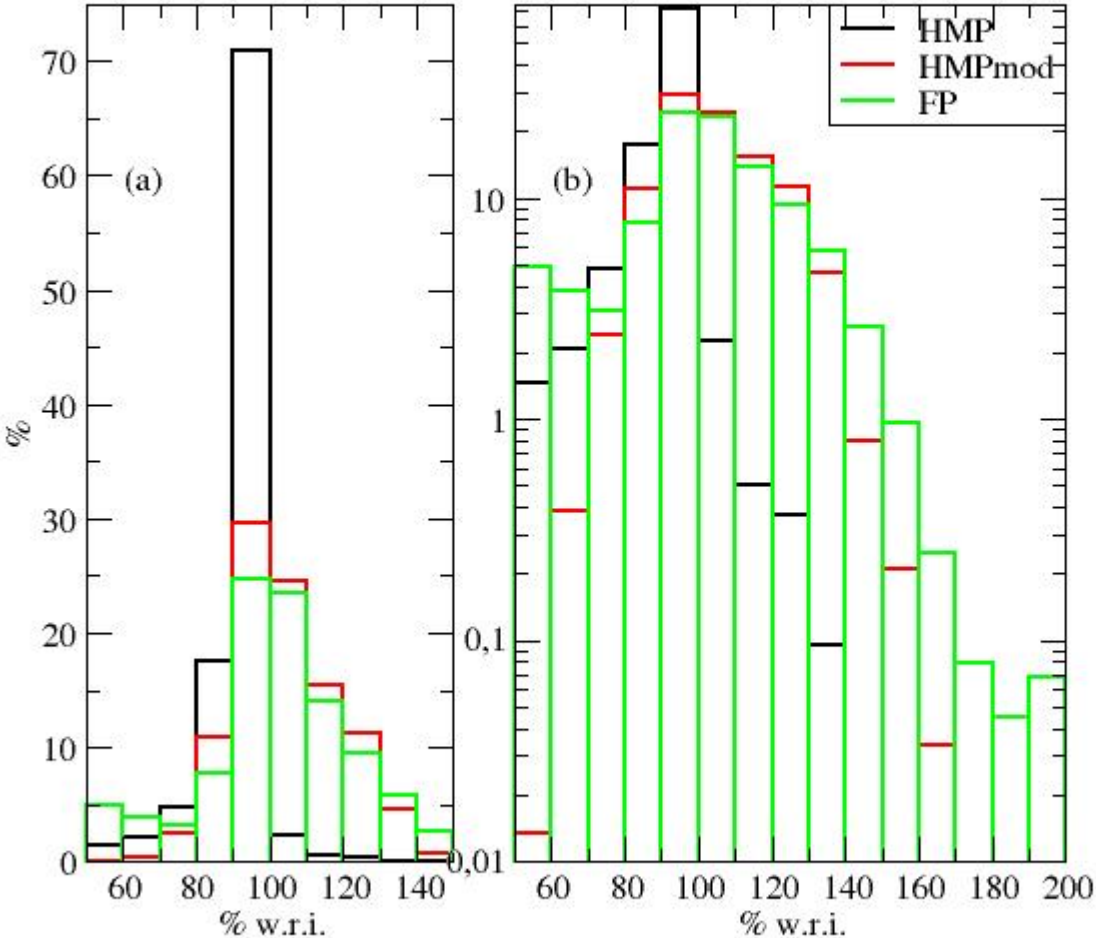
Challenge : how to measure moisture content of a supersaturated atmosphere?

Since any surface – including measuring device – acts as IN?



- Simple (use commercial sensors)
- Robust (deploy, operate in extreme environment)
- Affordable (can deploy multiple units , e.g. for vertical gradients)

Observation confirm that supersaturation is very frequent, the norm rather than the exception



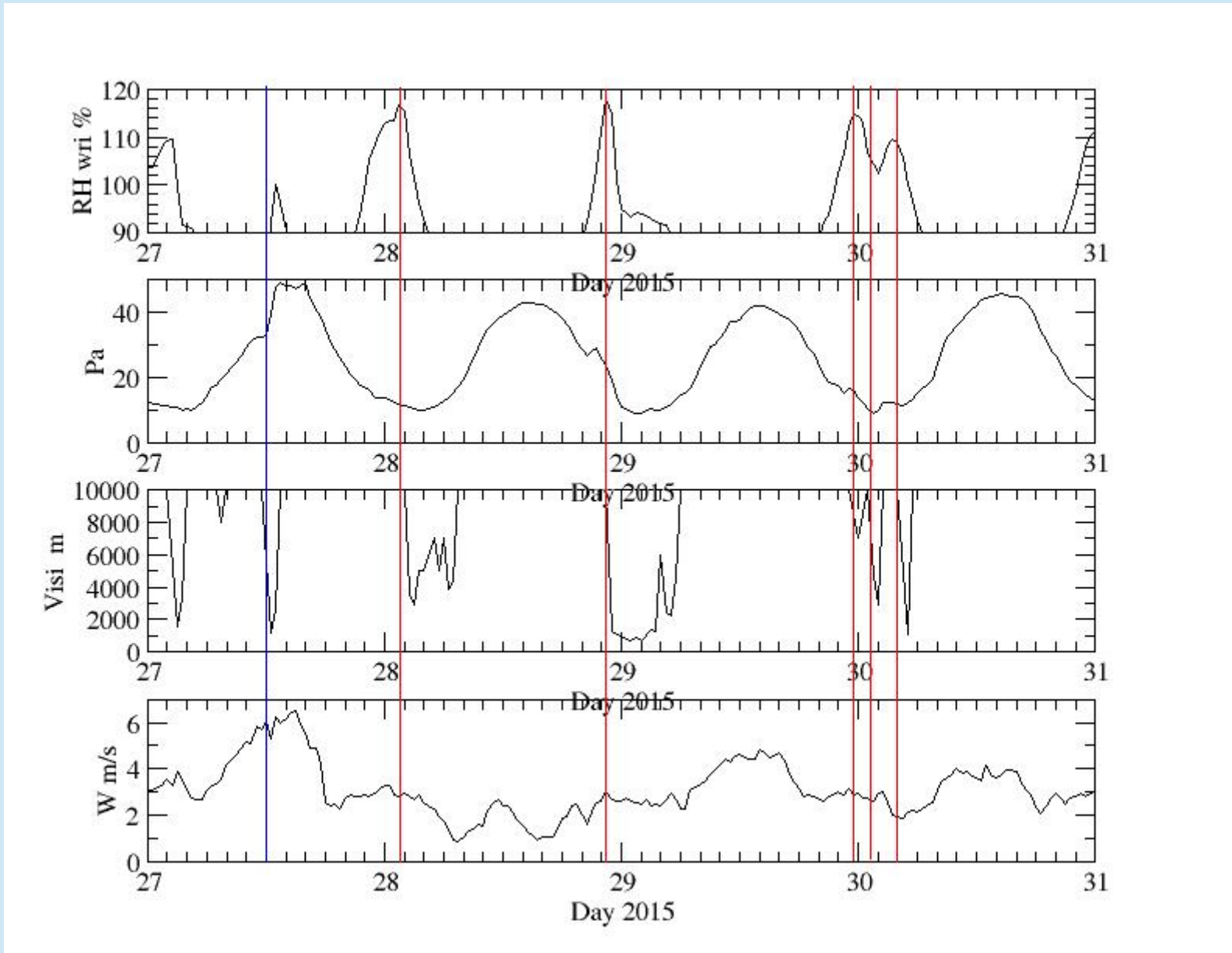
The Dome C surface atmosphere is characterized by

- Frequent supersaturations
- In conditions similar to high troposphere regarding temperature, moisture content, CCN and IN content
- But comparatively more accessible for in situ measurement (e.g. long continuous series)

Could it be used as a natural laboratory to test / evaluate / develop cold microphysics parametrizations of cold clouds and precipitation ?

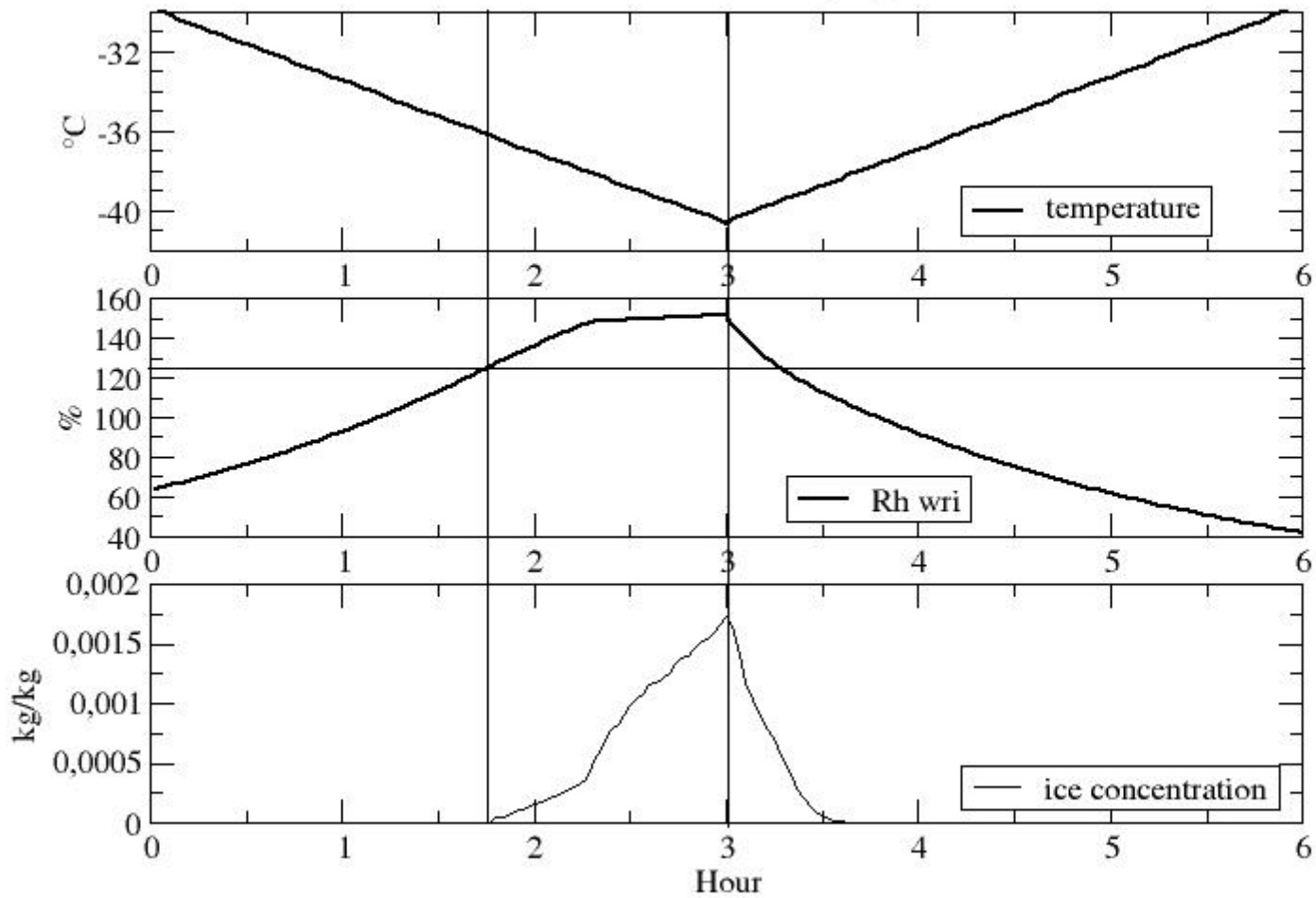


# Visibility from METAR (Meteorological Airport Reports) : diurnal haze formation in summer



# Microphysics Thompson 09

Kinematic Driver (KiD) model Shipway and Hill 2012

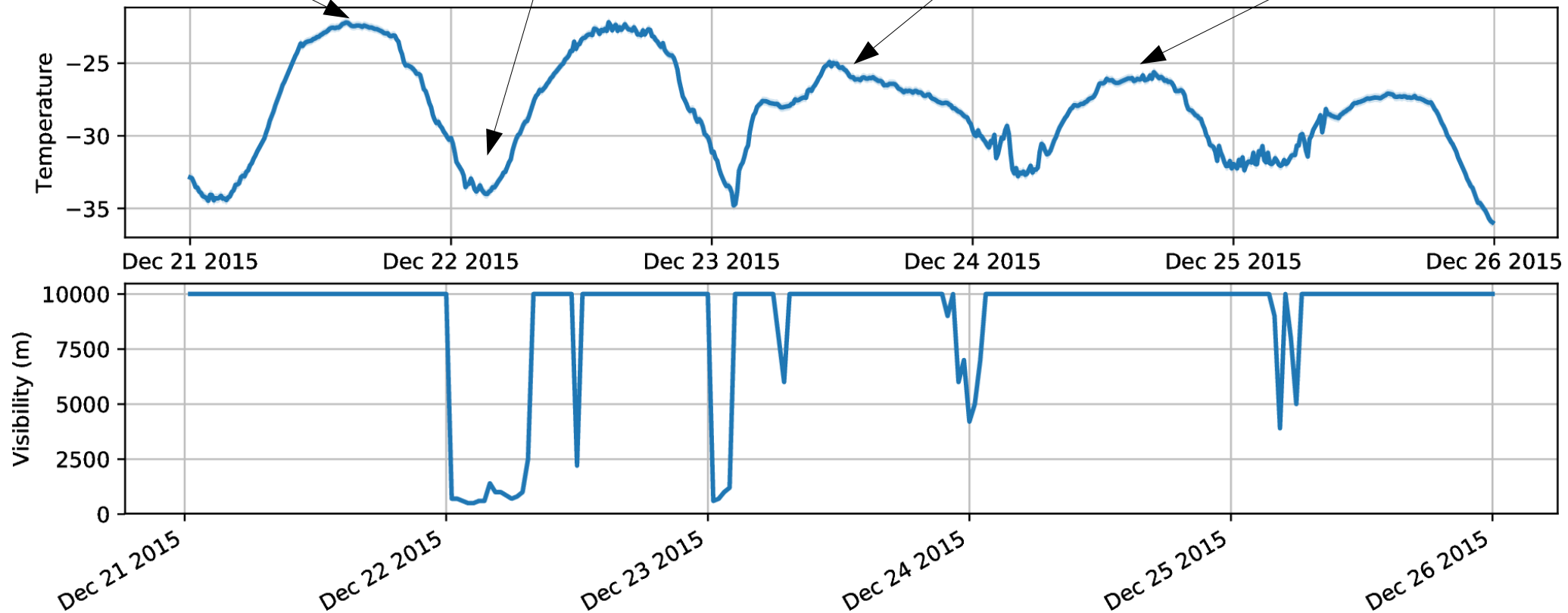




## A model evaluation / intercomparison (GAWS?)

J.-B. Madeleine, With R. Forbes, H. Gallée, E. Vignon, F. Hourdin, F. Lemonnier

# Selected period – Dome C obs.









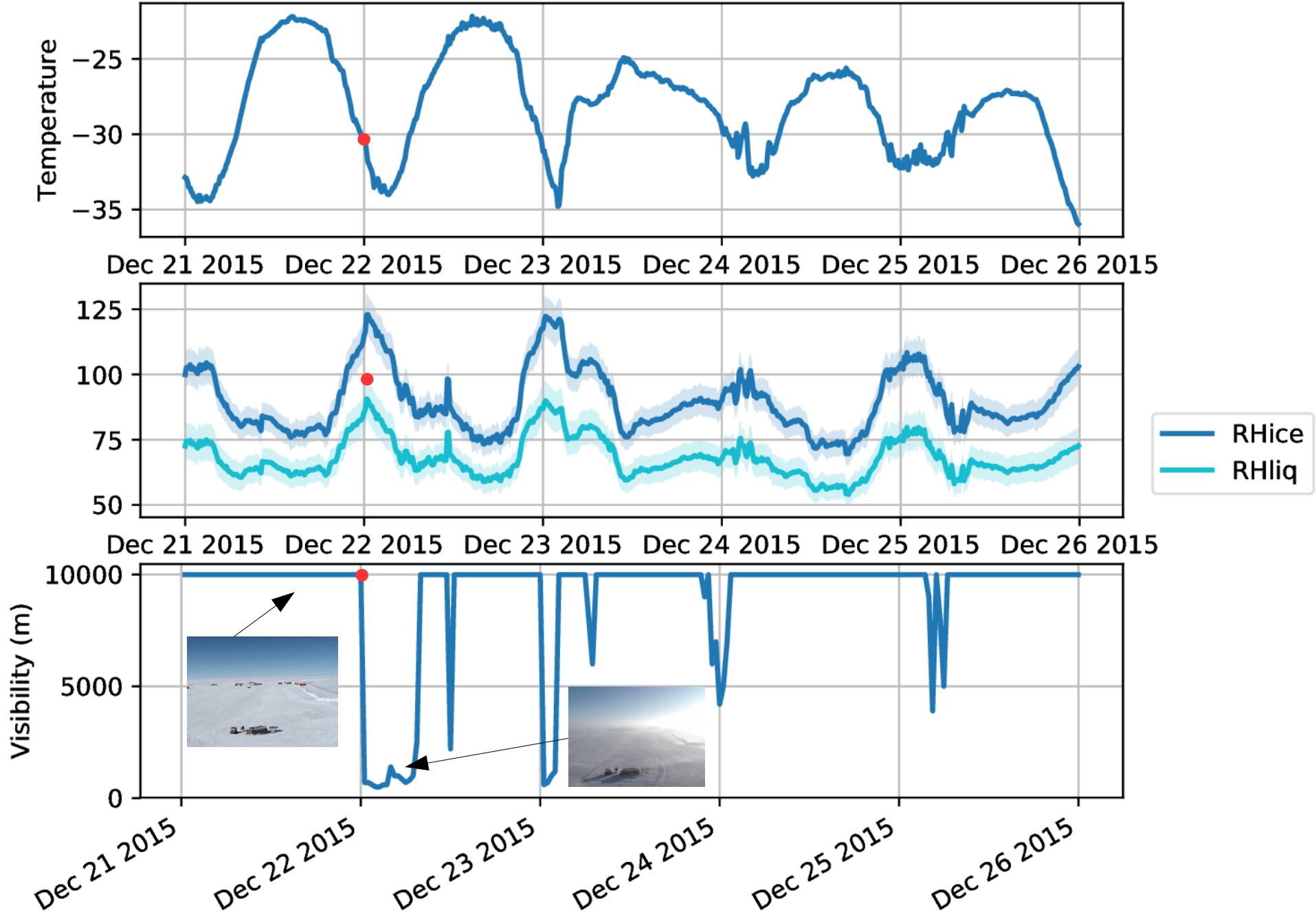








# Ice haze formation

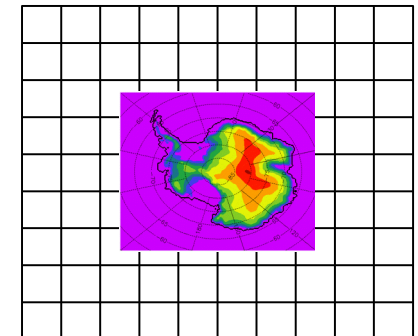
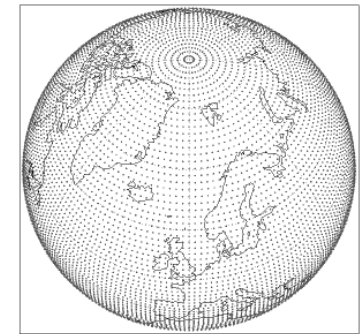
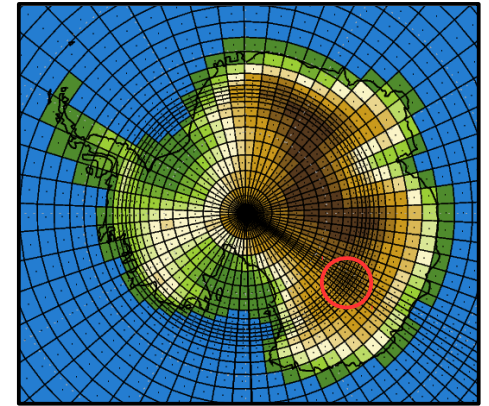


# Model intercomparison

- **IPSL-CM** (global, zoom) : no supersaturation  $\rightarrow$  ice condenses when  $RH_i = 100\%$

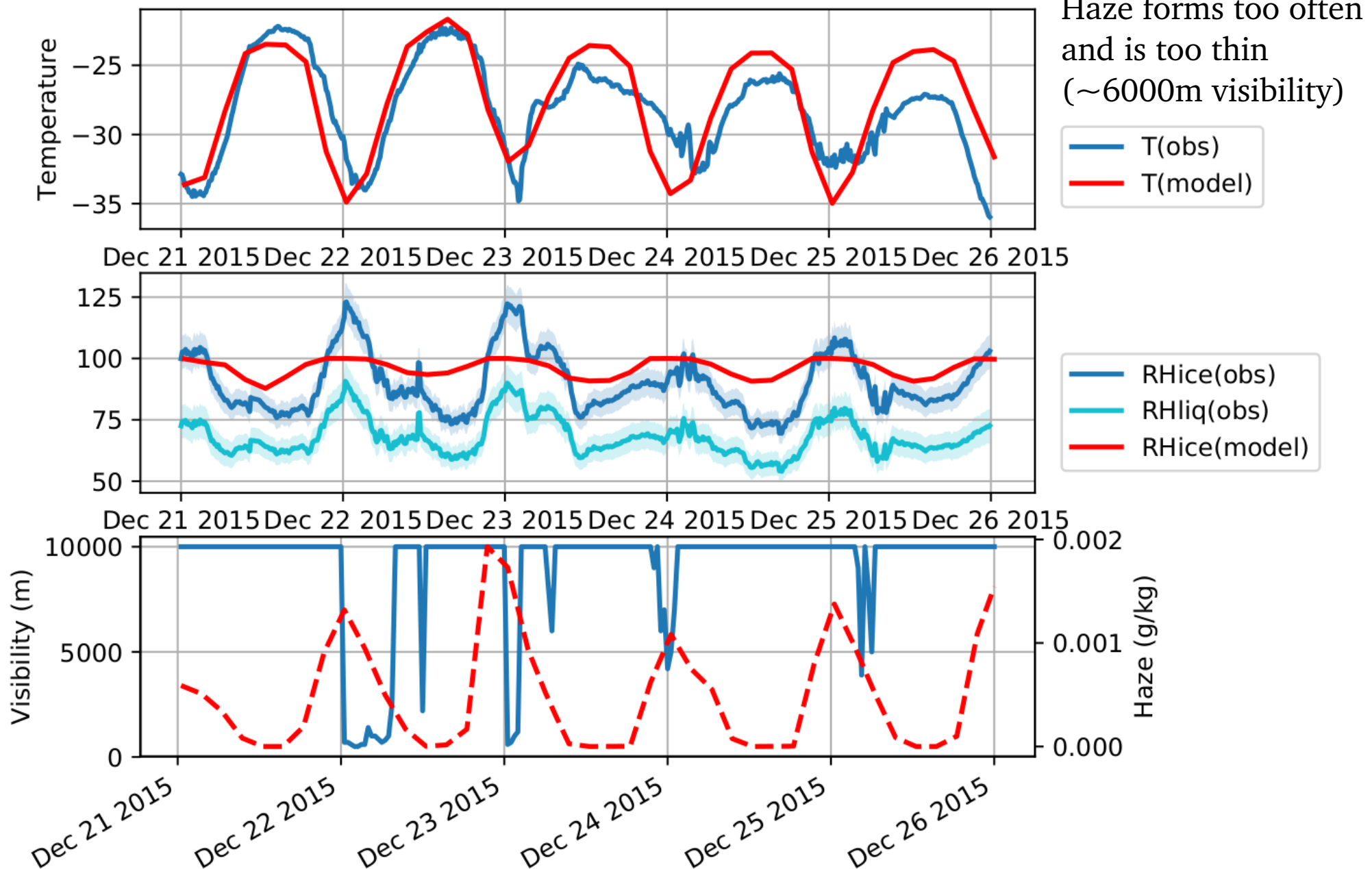
and Cloud liquid fraction =  $\left( \frac{T - T_{\min}}{T_{\max} - T_{\min}} \right)^n$

- **IFS** (global, high resolution) : microphysical scheme  $\rightarrow$  crystal growth by deposition for  $T > -38^\circ\text{C}$  and homogeneous freezing for  $T < -38^\circ\text{C}$
- **MAR** (limited area model) : microphysical scheme and blowing snow, 6 prognostic equations



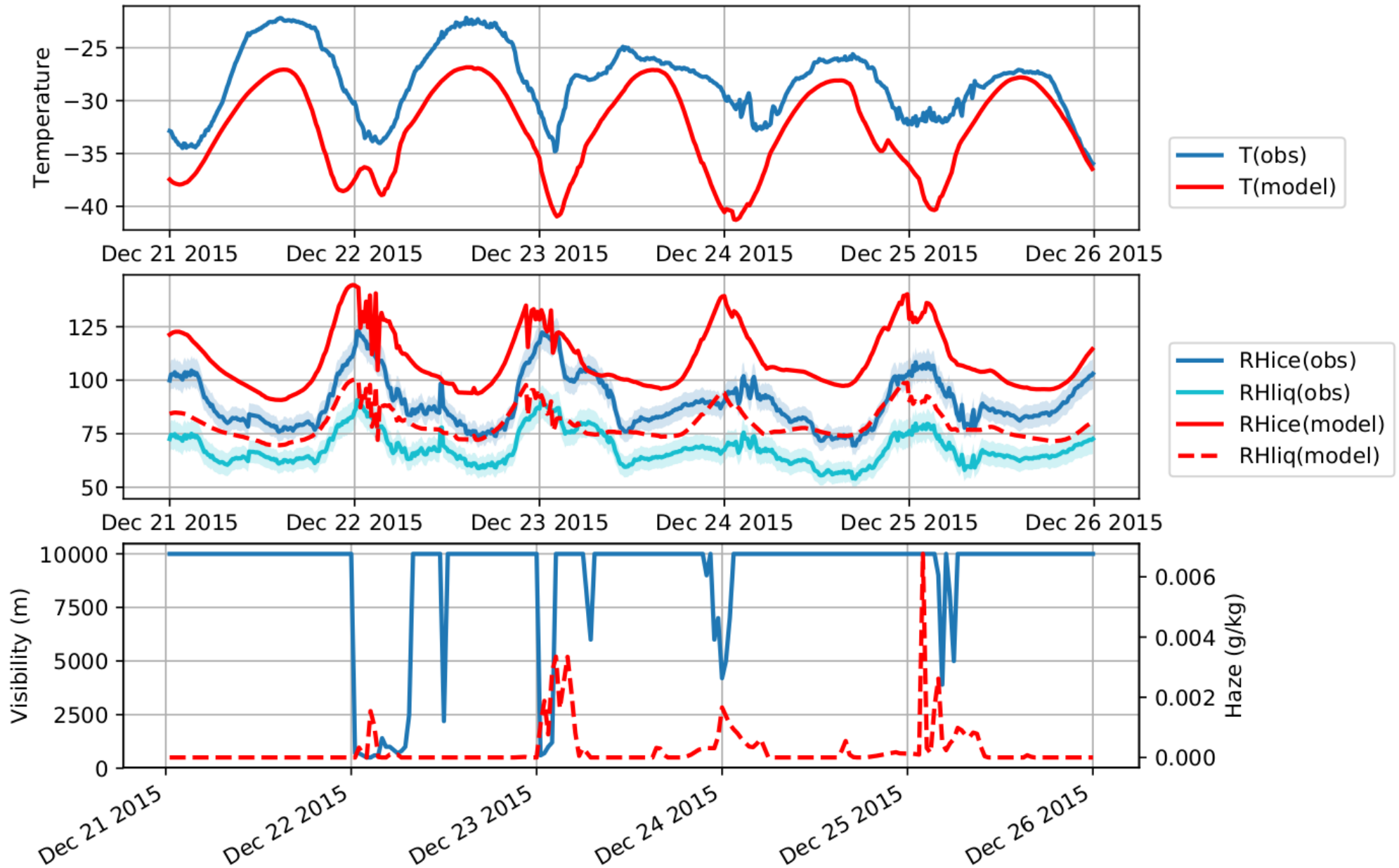
# Results without supersaturation

LMDz model results

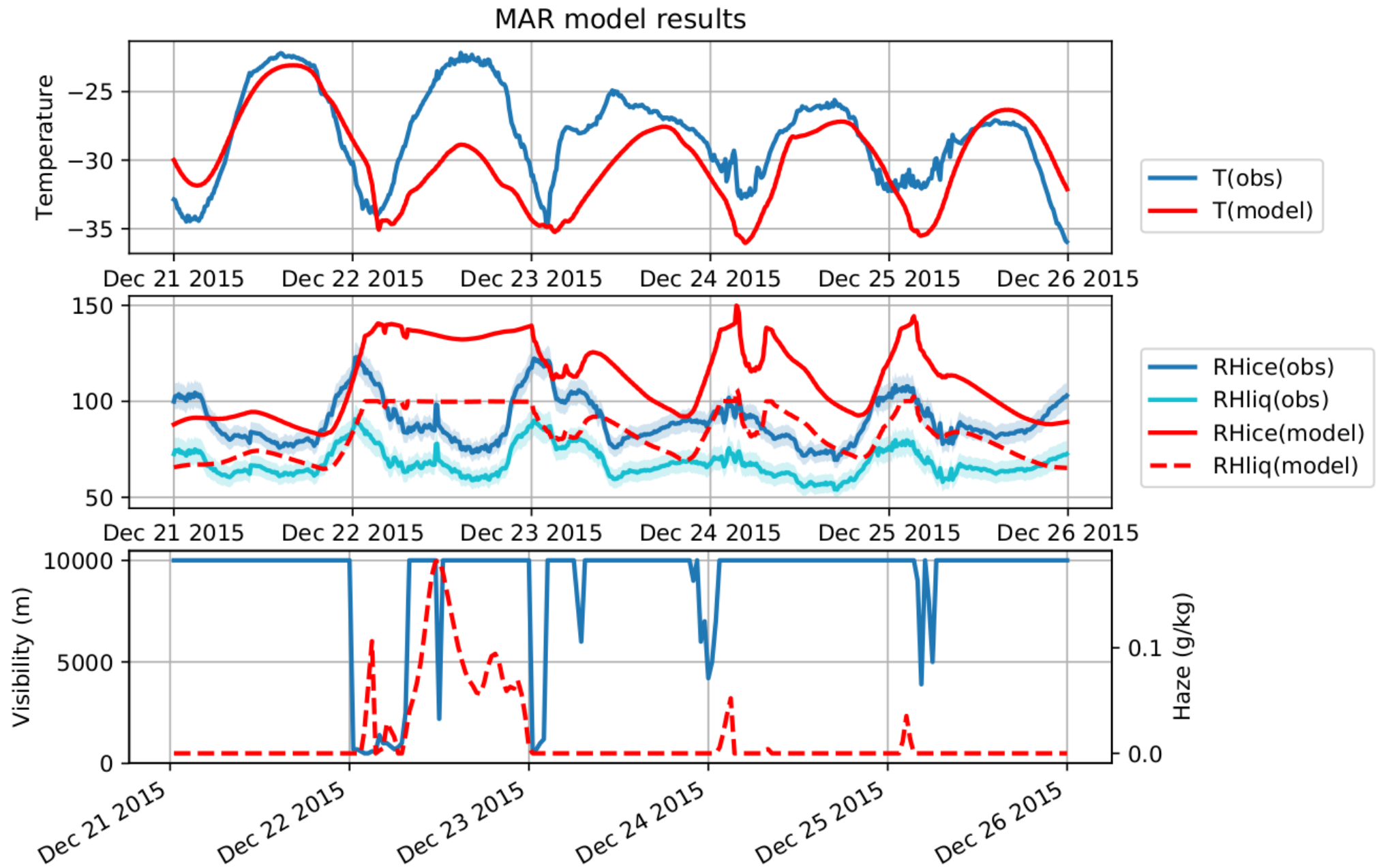


# IFS results (includes supersaturation)

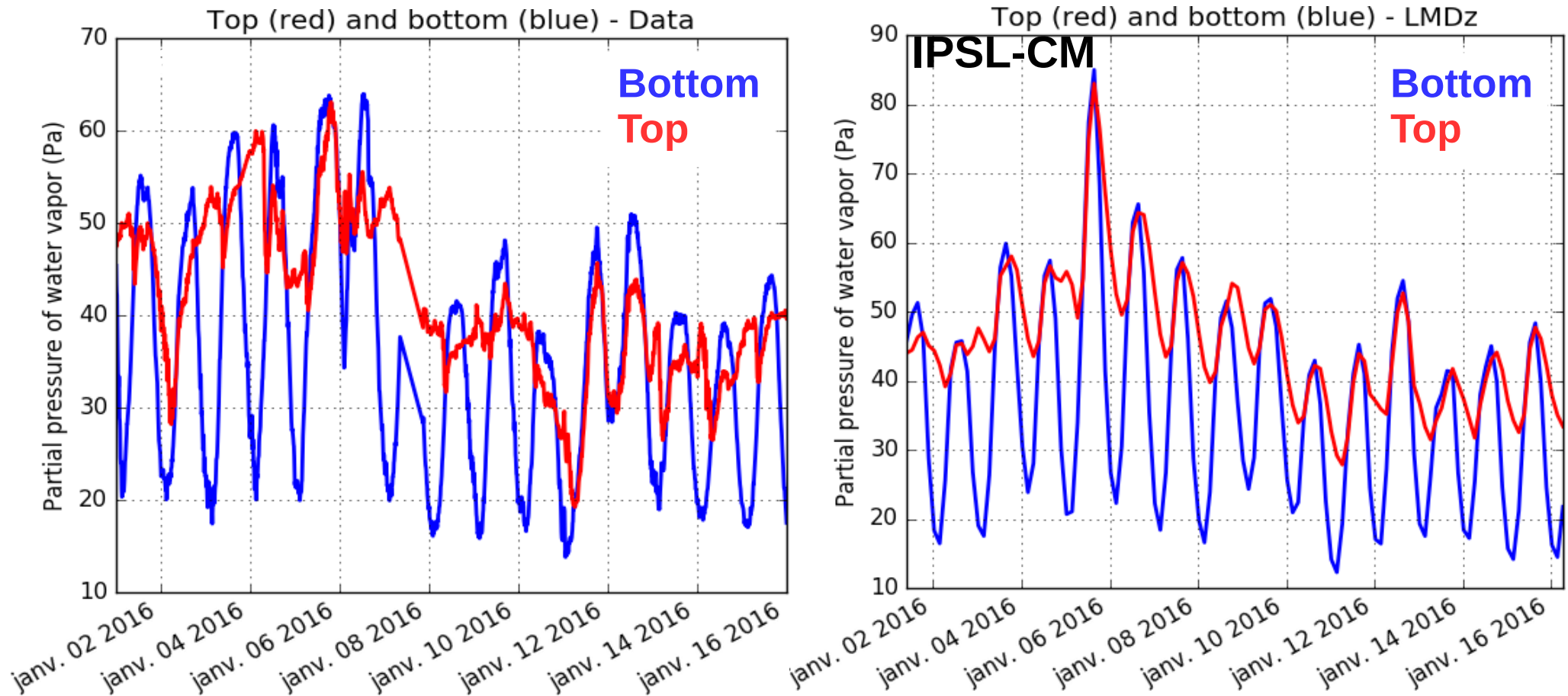
IFS model results (ECMWF)



# MAR results (40 km resolution)



# Vapor partial pressure along the mast



Turbulent diffusion in the boundary layer can be validated using the multiple hygrometers along the mast (gives access to the water vapor

# Conclusions

- IPSL-CM temperatures are in good agreement with the observations if albedo / heat capacity tuned, but the model underestimates haze opacity
- IFS shows good agreement with both the observed humidities and haze opacities despite a cold bias
- MAR overestimates supersaturation and haze opacity → fails to dissipate the haze → cold bias

These microphysical observations at Dome C are very promising  
→ work is underway to further constrain microphysics  
(amount of ice nuclei, microphysical properties of the fog)



**THANK YOU**